

The top of the martian atmosphere

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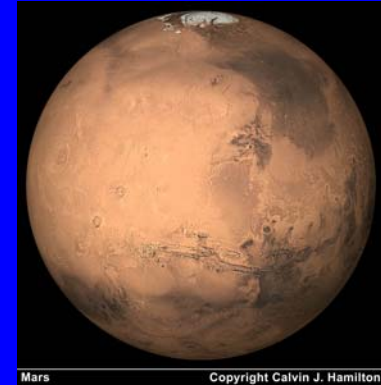
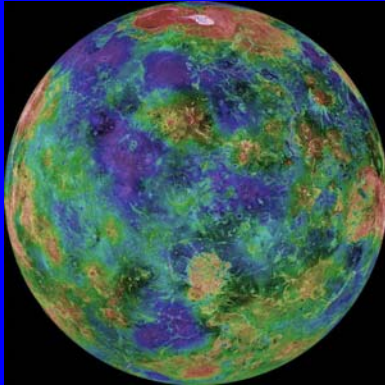
Images

- www.solarviews.com
- The Cosmic Perspective, Bennett et al.

Aim: Describe the current state of the Mars upper atmosphere and its important processes

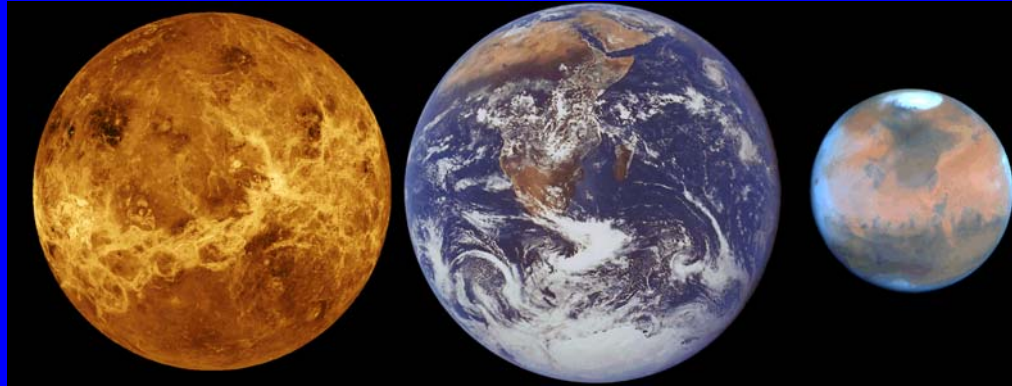
- Motivation: Atmospheric loss on Venus, Earth, and Mars
- Context: Present-day Venus, Earth, and Mars
- Details: Chemistry, dynamics, energetics, and plasma in the Mars upper atmosphere

Venus, Earth, and Mars – Some Similarities



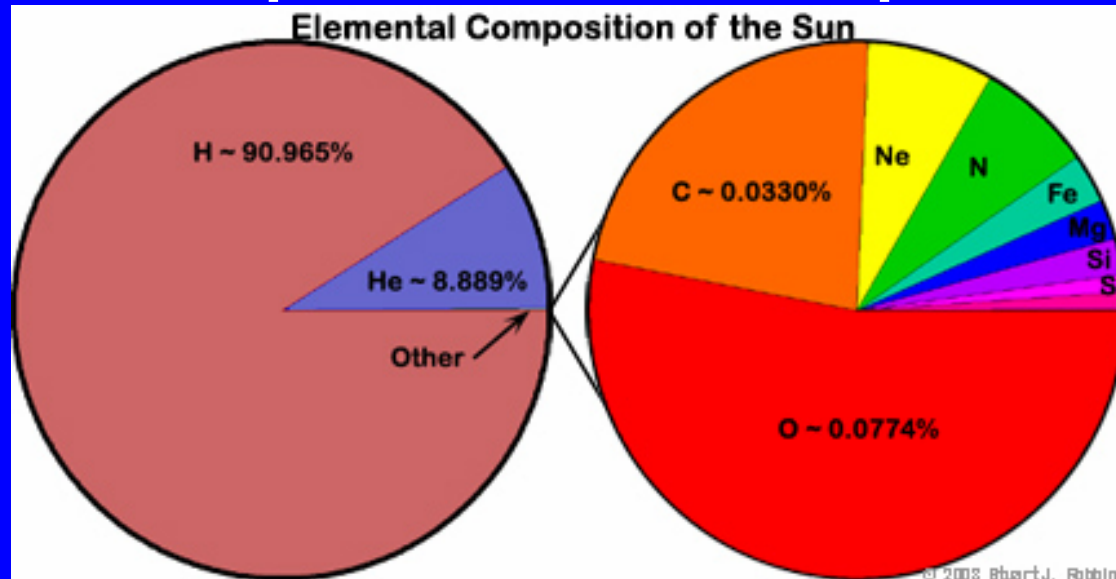
- Rock/metal surfaces and interiors
- Similar bulk compositions, so probably similar primordial atmospheres
- Substantial atmospheres today
- Similar sizes and gravities

Venus, Earth, and Mars – Some Differences



- Distance from Sun: Mars > Earth > Venus, so expect Venus to be hot and Mars cold
- Magnetic field: Earth has strong dipole field, Venus has no field, Mars has regions of strong field and regions of no field
- Size: Mars is smaller than Earth and Venus, so its gravity is weaker

What gases should be common in terrestrial planet atmospheres?



- H_2O ?
- CO_2 ?
- N_2 ?

• Noble gases ?

• O_2 ?

• Must be abundant in the inner solar system

• Must be volatile enough to be gaseous

http://filer.case.edu/~sjr16/media/sun_elements.jpg

Atmospheric Compositions

- Meteorites, Sun, giant planets suggest that H_2O , CO_2 , N_2 should be most abundant
- Venus: 100 bar pressure
 - Mostly CO_2 , some N_2
- Earth: 1 bar pressure
 - N_2 and O_2 mixture
- Mars: 0.006 bar pressure
 - Mostly CO_2 , some N_2

Atmospheric History

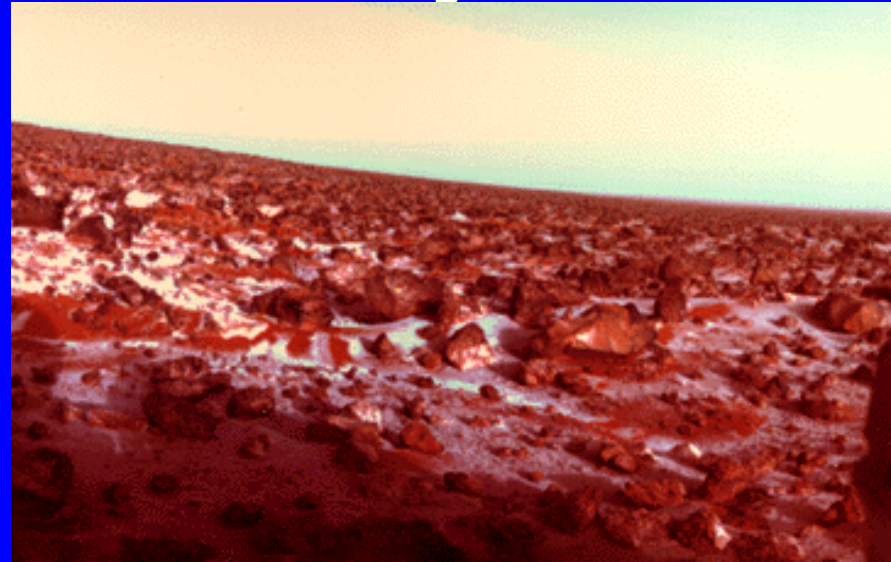
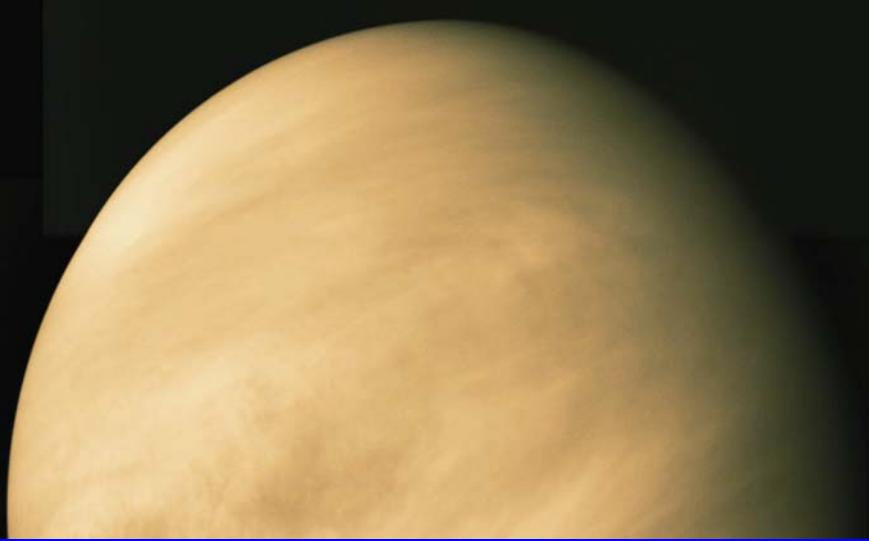


http://www.beachyhead.org.uk/_images/_images/_new/7sisters.jpg

http://www.aerofun.ch/image_gallery2/image9.jpg

- Earth: H_2O forms oceans, cannot be dissociated by UV, so no H lost to space. CO_2 dissolves in oceans, forms carbonate rocks. Leaves N_2 dominant gas in atmosphere. Continuous biological activity pumps O_2 into atmosphere and chemical reactions consume O_2

Atmospheric History



<http://www-curator.jsc.nasa.gov/antmet/marsmets/images/viking.gif>

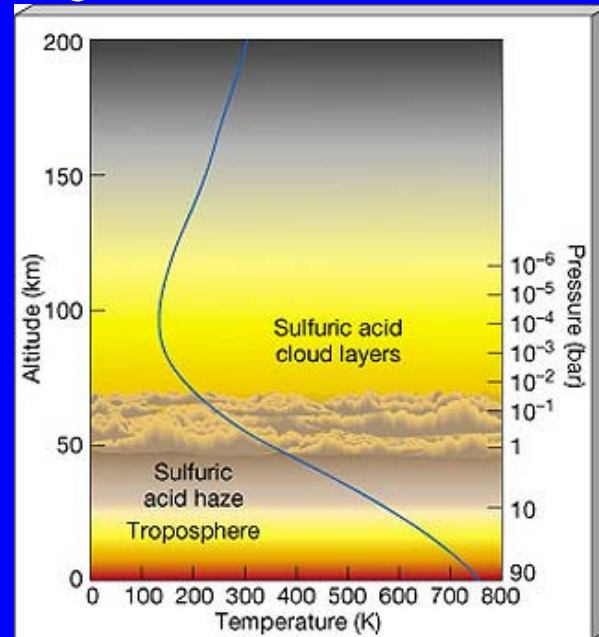
- Venus: Too Hot
 - H_2O in atmosphere photodissociates, H lost to space very rapidly. C, N, O lost more slowly, so CO_2 and N_2 remain in atmosphere
- Mars: Too cold
 - H_2O frozen solid. Low gravity and weak magnetic field mean that CO_2 and N_2 escape rapidly, reducing atmospheric thickness. Some frozen H_2O sublimates into thin atmosphere, is photodissociated and H is lost to space.
- The masses and chemical compositions are explained! But this “story” needs rigorous testing.

Escape Questions

- How did Mars lose 99.9% of primordial atmosphere, yet keep 0.1%?
- What is the history of Mars climate?
- How have magnetic fields affected escape?
- What are dominant escape processes today and how do they operate?
- What is upper atmosphere like today and what processes are important?

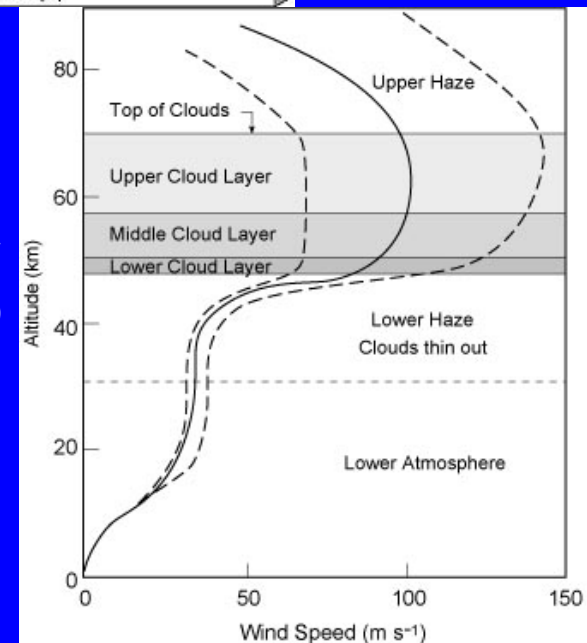
Present-Day Venus

- Zero obliquity and eccentricity mean no seasons
- Thick atmosphere, slow rotation mean that weather at surface is same everywhere
- 740 K at surface, slow winds, no storms, no rain
- H_2SO_4 clouds at 50 km, where pressures and temperatures are similar to Earth's surface



<http://rst.gsfc.nasa.gov/Section19/Venus-atmos-profile-CM.jpg>

http://as.e.tufts.edu/cosmos/pictures/Explore_figs_8/Chapter7/Fig7_3copy.jpg

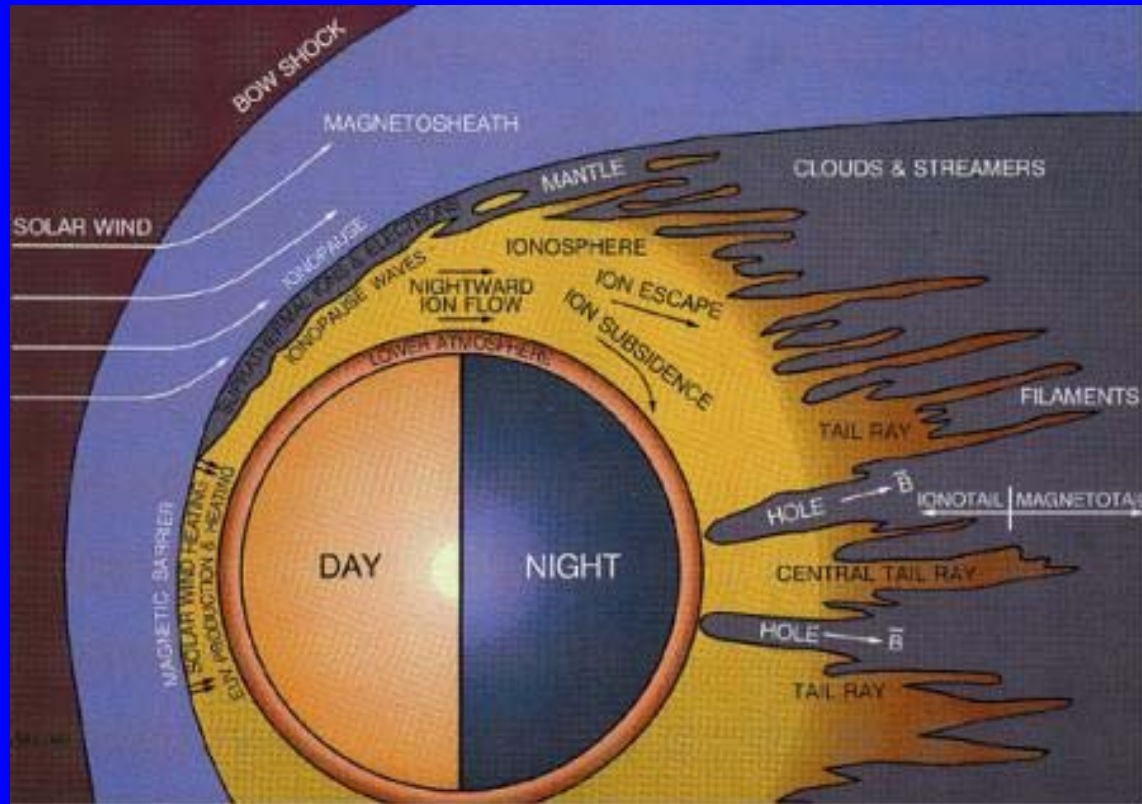


Venus upper atmosphere

- Reformation of photolysed CO_2 catalysed by Cl, terrestrial implications
- Lots of solar heating, but little day-night transport of energy
- Nightside upper atmosphere is very cold, 100 K, whereas dayside is 300 K
- O / CO_2 ratio plays a major role, more O than CO_2 above 150 km
- Only H is escaping today

Venus ionosphere and plasma

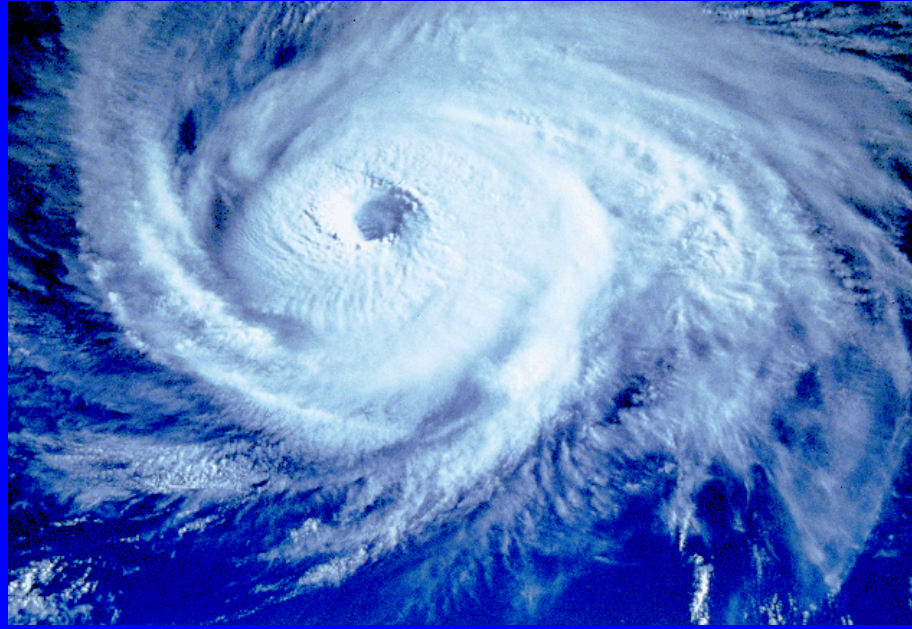
- Ionosphere formed by EUV photoionization of CO_2 , but $\text{CO}_2^+ + \text{O} \rightarrow \text{O}_2^+ + \text{CO}$
- O_2^+ is dominant at Chapman peak (140 km), O^+ dominant 40 km higher up



http://www3.imperial.ac.uk/spat/research/space_magnetometer_laboratory/spacemissionpages/venusexpresshomepage/science

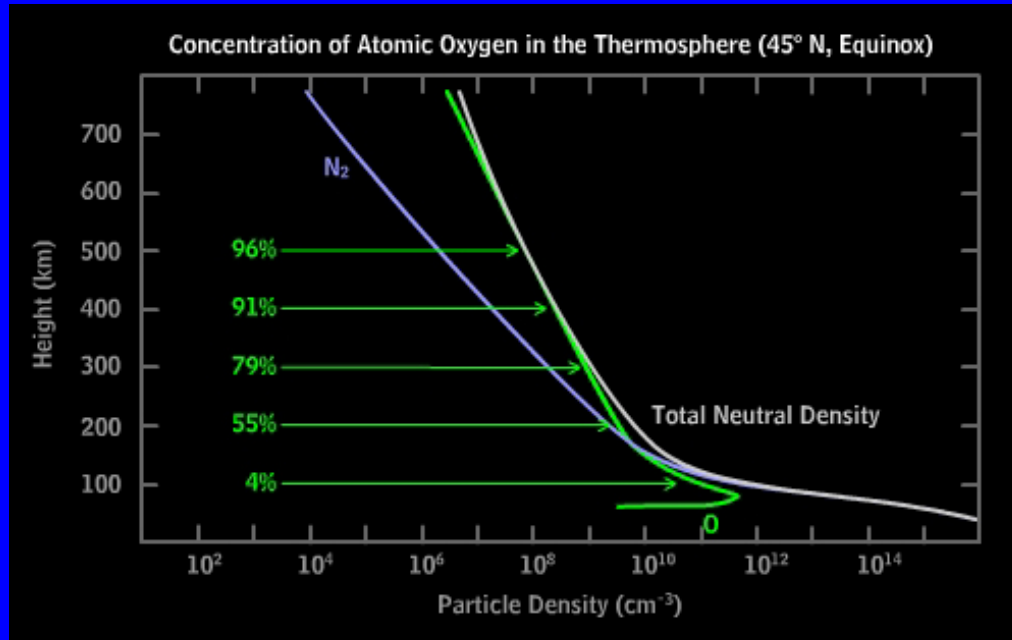
- Transport important near O_2^+/O^+ transition and above
- Magnetic fields due to draping of solar wind around planet
- Nightside ionosphere and magnetic fields are complex and variable, affected by plasma transport across terminator

Present-Day Earth



- Coupled atmosphere and ocean
- Strong seasonal and latitudinal variations, but significant transport of heat from dayside to nightside and equator to poles
- Rapid rotation and ocean/land contrasts drive lots of weather
- Stable global circulation patterns (Hadley cells)

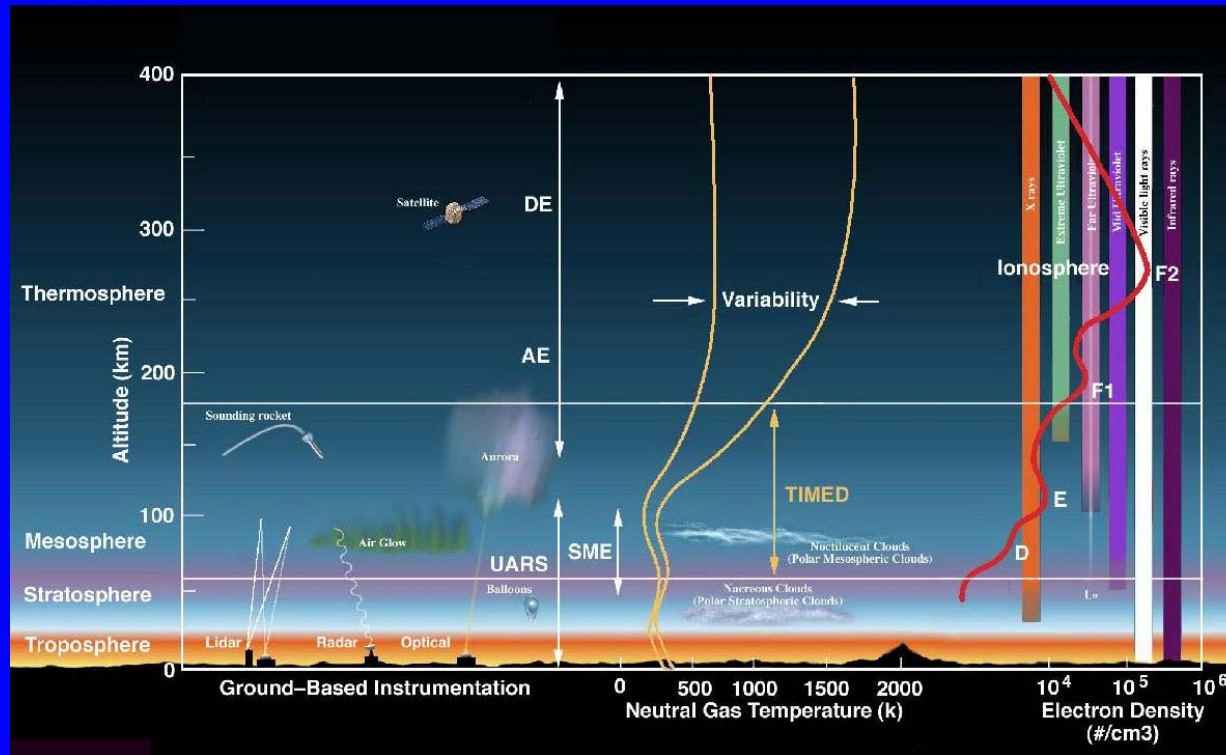
Earth Upper Atmosphere



http://www.meted.ucar.edu/hao/aurora/images/o_concentration.jpg

- O is more abundant than O_2 above 100 km and more abundant than N_2 above 200 km
- $T > 800 \text{ K}$ above 200 km, much hotter than Venus or Mars. These atmospheres are cooler because CO_2 is very effective at radiating heat, whereas Earth needs higher temperature gradients to conduct heat downwards
- Heating at poles due to magnetic fields guiding solar wind
- Only H is escaping today

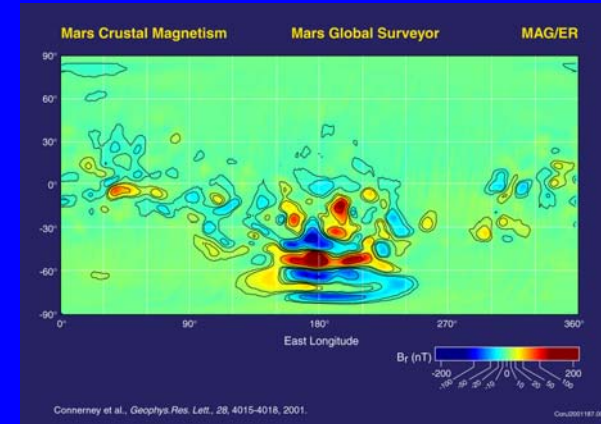
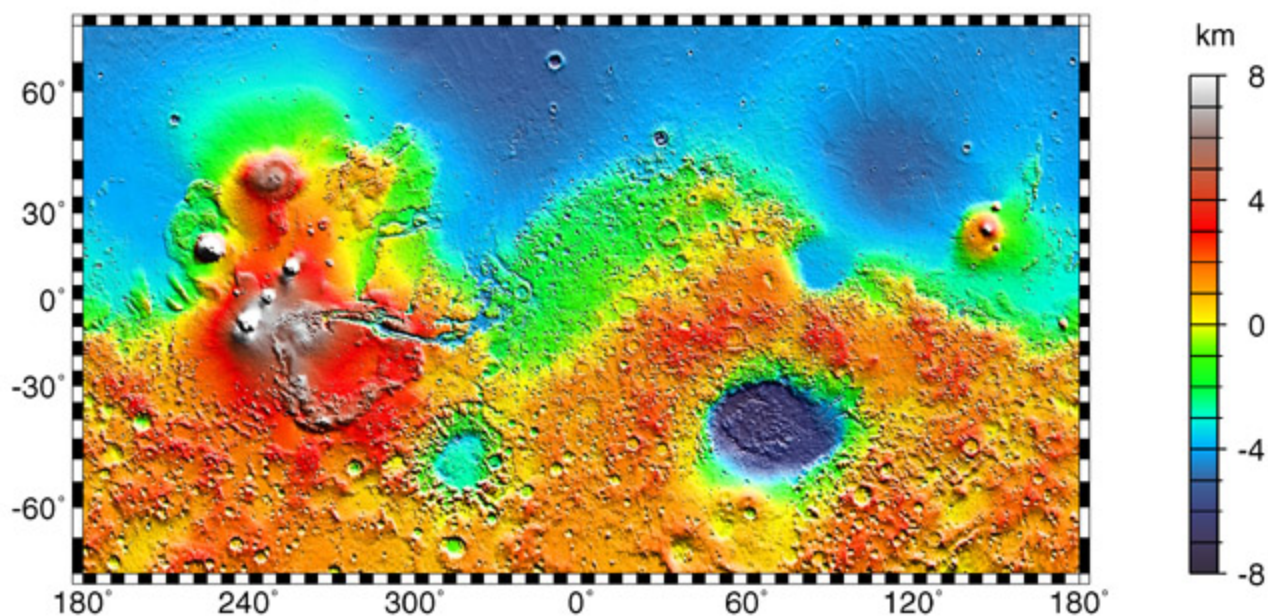
Earth Ionosphere



http://www.bu.edu/cism/CISM_Thrusts/ITM.jpg

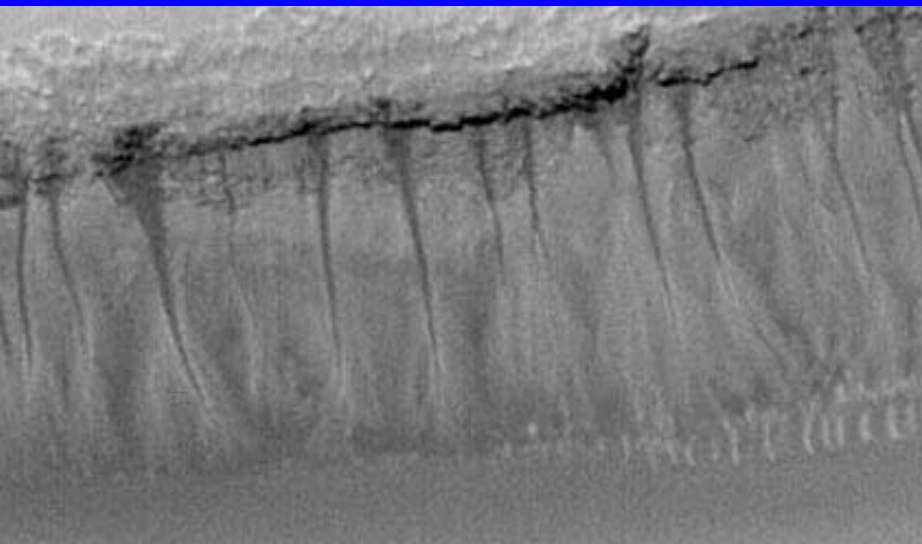
- O_2^+ and NO^+ dominant at 100 km, where EUV absorption peaks
- O^+ dominant at 300 km (overall peak), where transport plays major role
- Changes from O_2^+/NO^+ to O^+ and from N_2 to O make things complex
- Magnetic fields affect plasma transport, especially at equator and near poles

Present-Day Mars



- Minimal wind erosion
- Possible weak ongoing volcanism and tectonism, no evidence for past plate tectonics
- Frozen H₂O at poles and elsewhere
- Frozen CO₂ at poles
- Large canyons, craters, volcanoes

Present-Day Mars



- 1/3 of atmosphere freezes onto winter polar cap
- Global dust storms
- Large day/night temperature differences
- Surface pressure too low for liquid water to be stable, but ongoing gully formation may require liquid water
- Saturated with H₂O, both H₂O and CO₂ clouds are common

Lower, Middle, Upper

- Lower: 0 – 40 km, exchanges heat with surface and/or suspended dust
- Middle: 40 – 100 km
- Upper: 100 - ?? km, heated by variable solar UV flux, temperature increases with increasing altitude
 - Contains homopause, ionosphere, exobase
- Unlike Earth, regions are not well-defined

The rest of the talk

- Describe upper atmospheric chemistry, dynamics, energetics, and ionosphere
 - Using one illustrative example for each
- Link these examples by how they might affect escape

Chemistry: O and CO₂

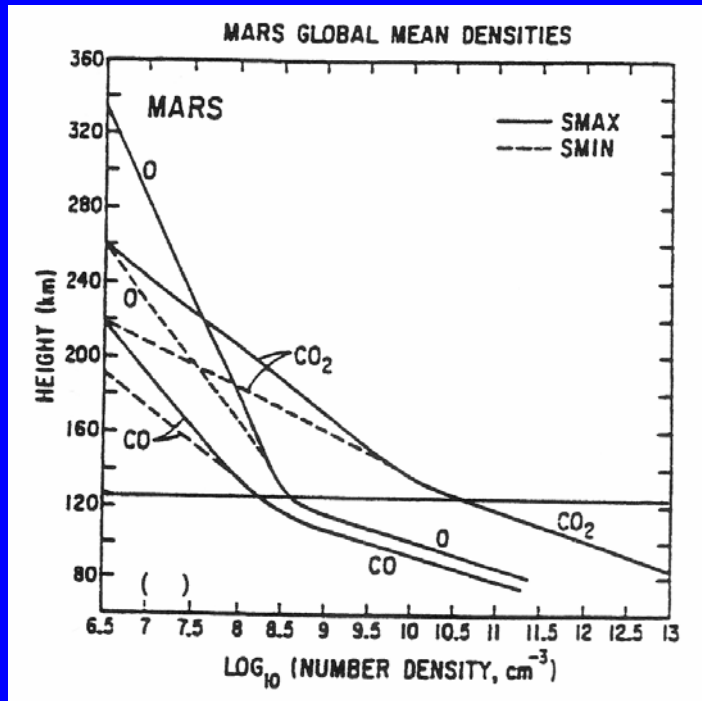
- CO₂ is photolysed by UV into CO and O
 - This is a rapid process
- $\text{CO}_2 + h\nu \rightarrow \text{CO} + \text{O} \ (\lambda < 1671 \text{ \AA})$
- $\text{O} + \text{O} + \text{CO}_2 \rightarrow \text{O}_2 + \text{CO}_2$

- Why is the atmosphere made of CO₂, not 2 parts CO and 1 part O₂?

Mars surface is very oxidizing

- Species like H, OH, HO₂, and H₂O₂ are relatively abundant and act as catalysts
- $\text{CO} + \text{OH} \rightarrow \text{CO}_2 + \text{H}$
- More chemistry oxidizes H back into OH
- Net effect is $\text{CO} + \text{O} \rightarrow \text{CO}_2$
- But not all O is lost...

O in the upper atmosphere

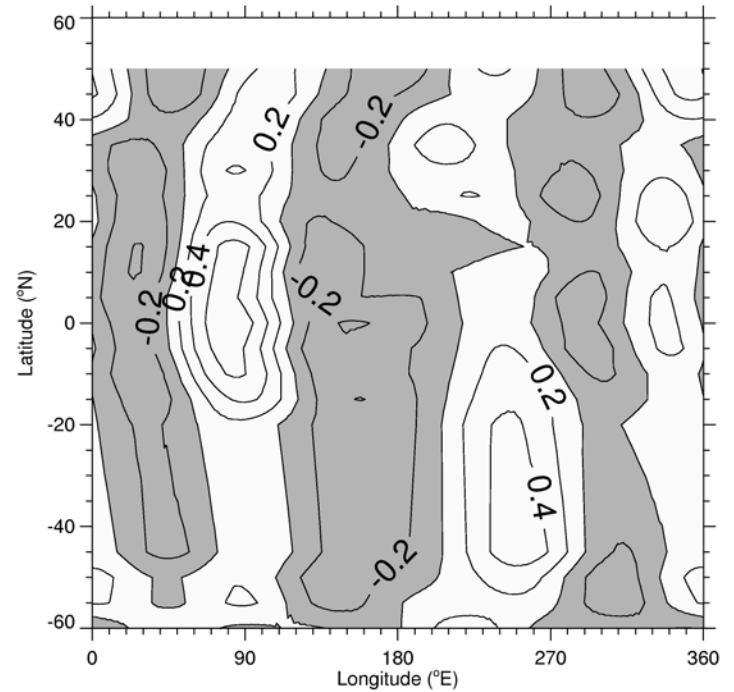
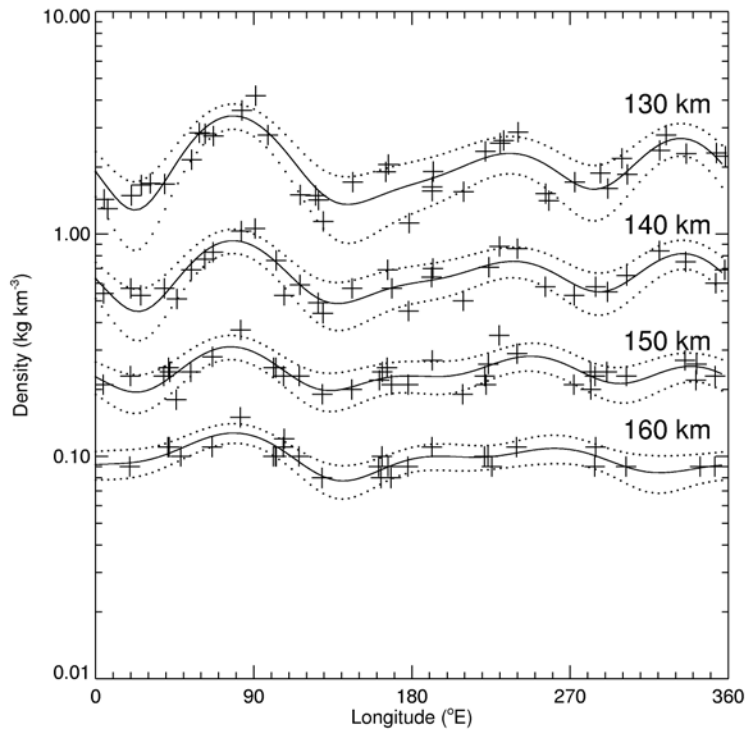


Bougher et al., 2002

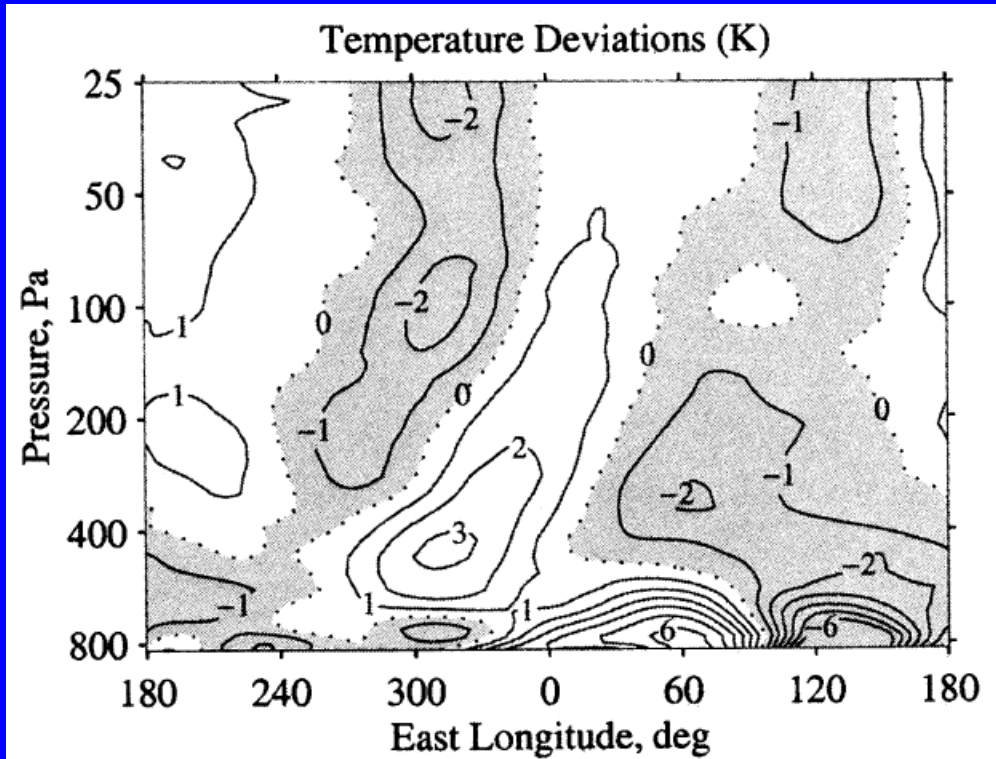
- $O / CO_2 > 1$ above 200 km
- O bumps into CO₂, excites vibrational states (demo)
- CO₂ emits 15 μ m photon, de-excites
- Important for cooling upper atmosphere

- EUV photons absorbed by CO₂, form CO₂⁺ ions
- But CO₂⁺ + O \rightarrow O₂⁺ + CO
- O makes O₂⁺ the dominant ion
- Presence of neutral O makes it easy for O to escape, affects temperature of upper atmosphere, and controls composition of the ionosphere

Dynamics - Tides



Source of tides



Hinson et al., 2001

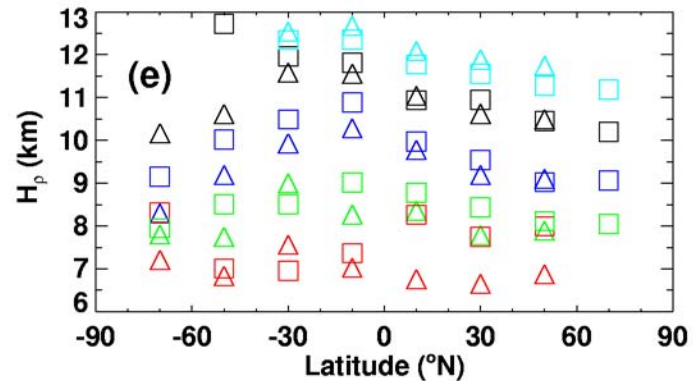
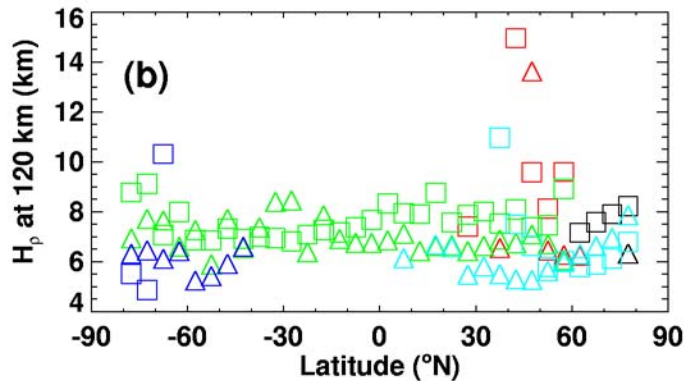
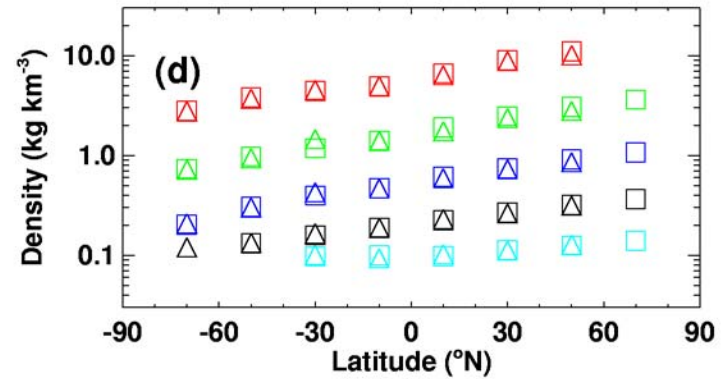
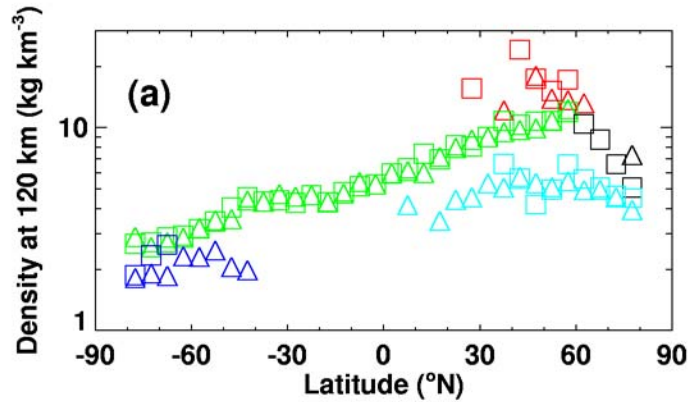
Some modes dissipate as they propagate up
Some modes amplify as they propagate up
Different modes dominant in lower and upper atmosphere

- Must be surface or interior
 - Nothing above surface varies with longitude
- Interaction of migrating tides with topography
- Dynamics of lower and upper atmospheres are linked together

Implications of tides

- Strong dynamics in upper atmosphere
- Circulation patterns will cause heating/cooling
- Breaking of waves and tides deposits energy and momentum into atmosphere
- Winds may transport plasma to regions where escape is easier

Energy – Temperature Variations



Panels a and b: Colours indicate different areobreaking seasons/LSTs
NP: Black = dayside, Light Blue = nightside
SP: Green = dayside, Dark Blue = nightside

Panels d and e: Colours indicate 120 – 160 km in 10 km intervals

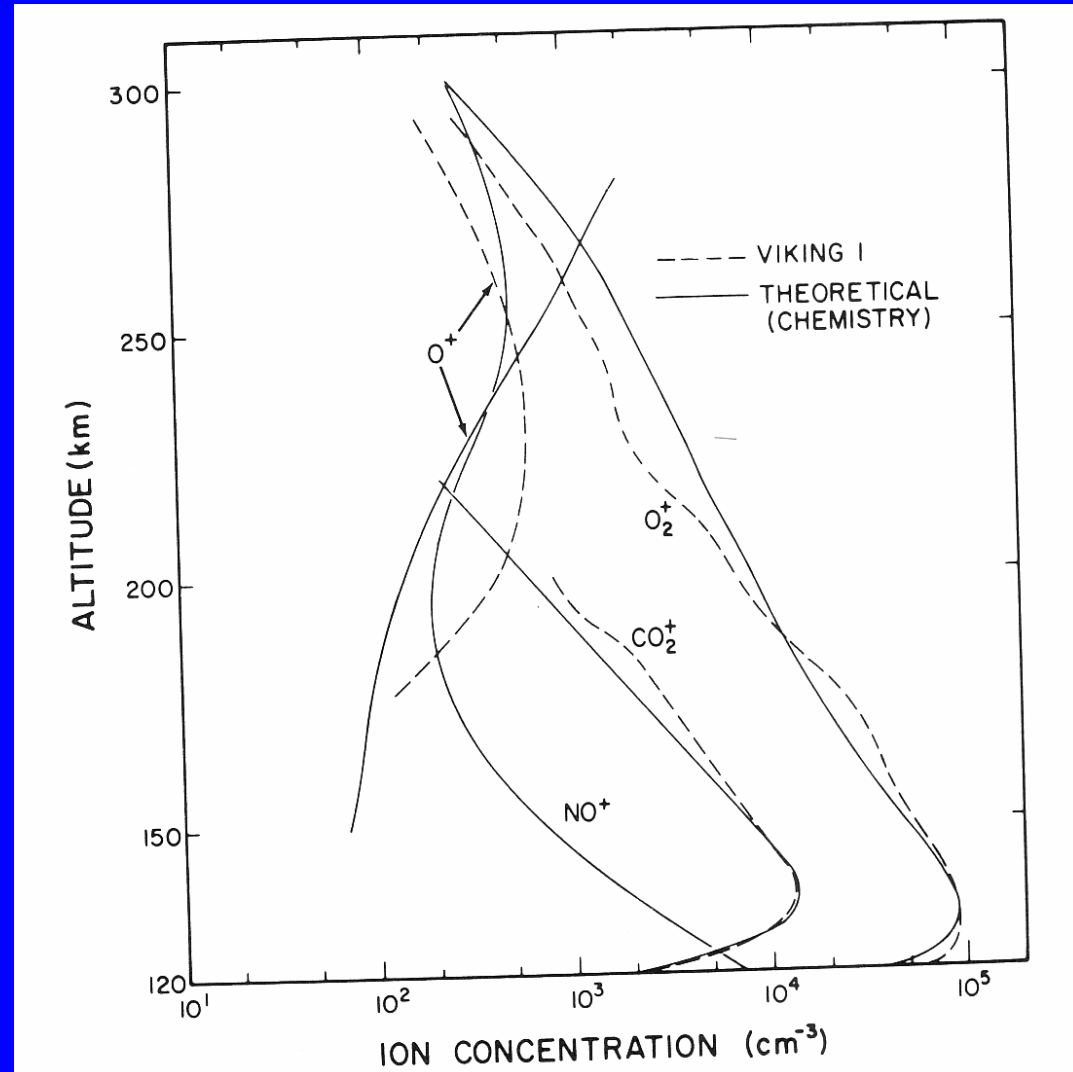
Temperature Observations

- Vertical temperature gradients much greater than meridional gradients
 - Atmospheric circulation transports heat poleward efficiently
 - Vertical gradients needed to conduct heat downwards into denser atmosphere where CO₂ 15 um radiation can radiate heat to space
- Mars: night T = 100K, day T = 150K
- Venus: night T = 100K, day T = 300K
 - Circulation is effective at heating Mars nightside
 - If too cold, nightside would not experience much escape

Ionosphere – Magnetic Fields

- $\text{CO}_2 + h\nu \rightarrow \text{CO}_2^+ + e$
- $\text{CO}_2^+ + \text{O} \rightarrow \text{O}_2^+ + \text{CO}$
- $\text{O}_2^+ + e \rightarrow \text{O} + \text{O}$

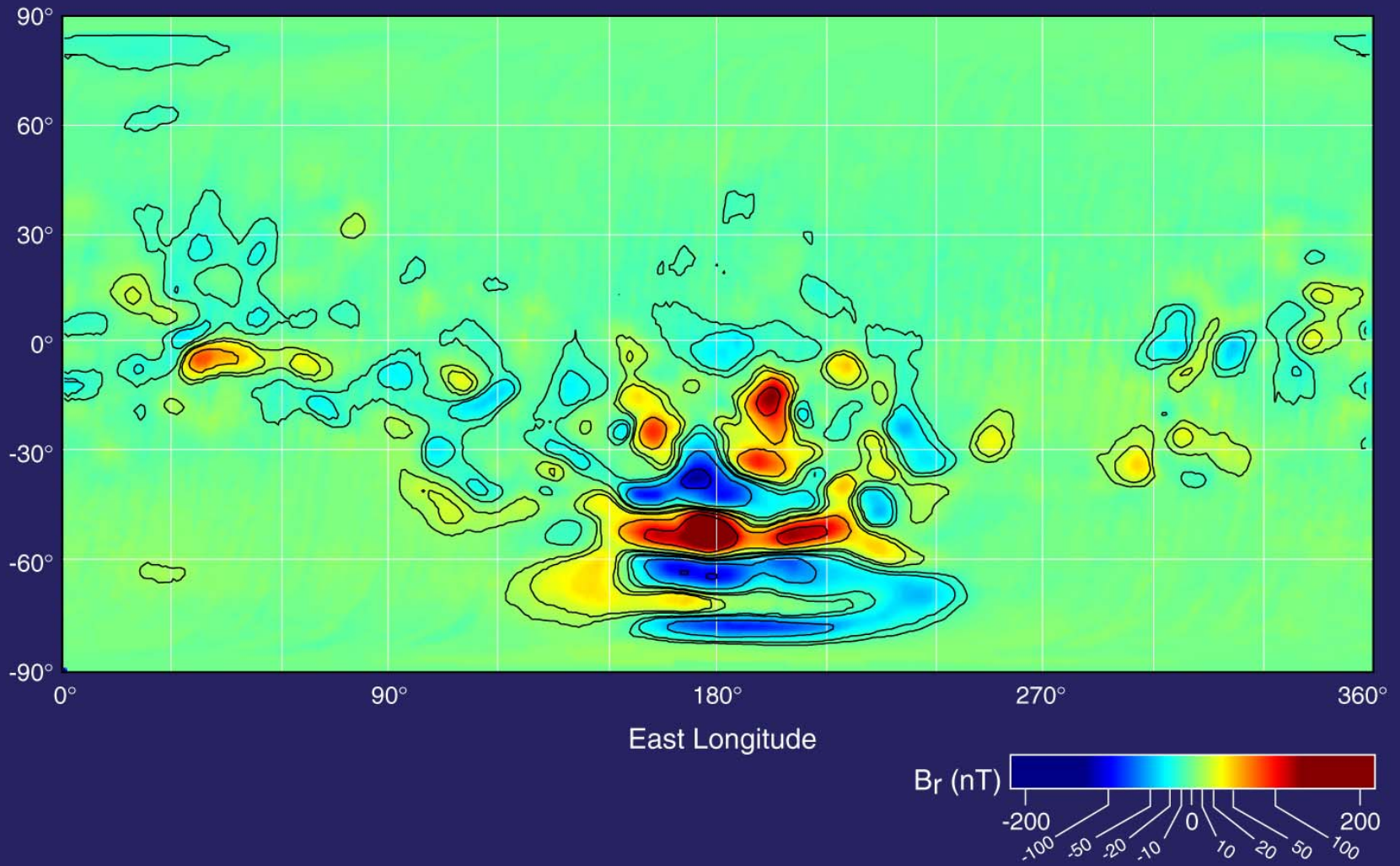
- Photochemical Chapman layer below ~ 180 km
- Transport important above there
- X-rays form second, lower peak that's hard to model
- Not enough ion or neutral composition data

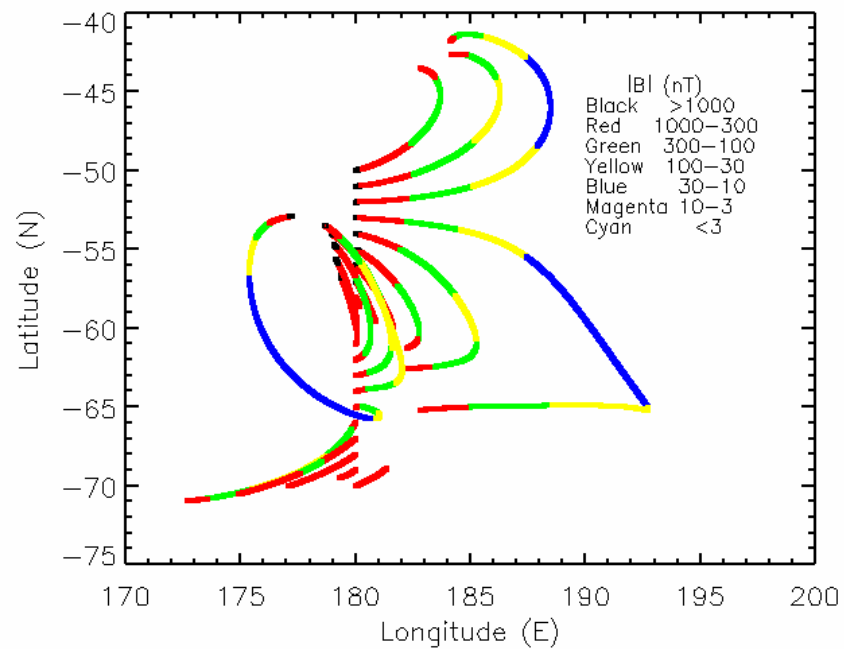


Mars Crustal Magnetism

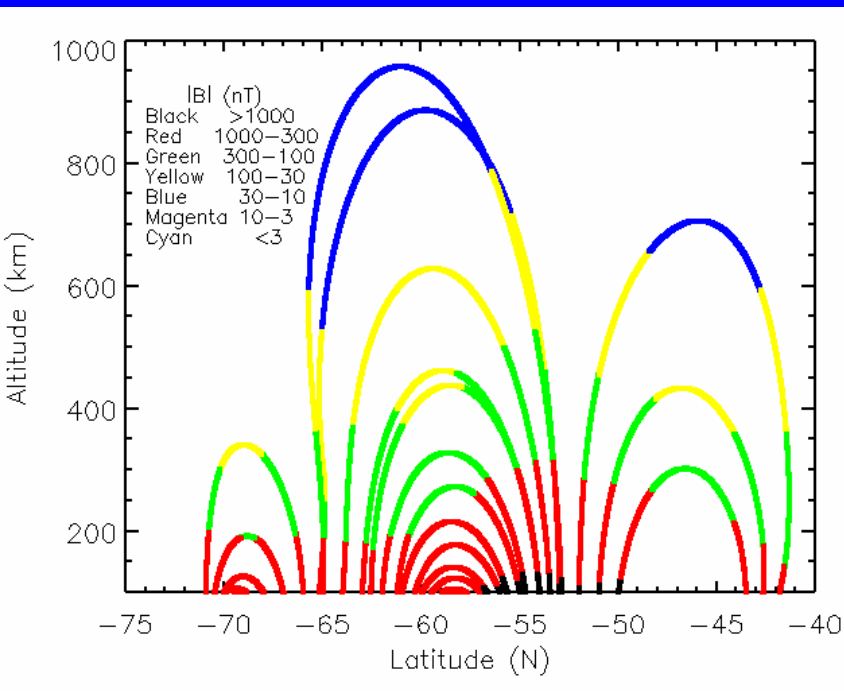
Mars Global Surveyor

MAG/ER





- This is not like Earth
- Field is not dipolar
- Field is not always strong
- Field geometry gives me a headache



- Plasma transport in Mars ionosphere is hard to model without magnetic fields
- How do magnetic fields affect transport?
- Next: I want to convince you that a problem exists. I do not offer a solution to the problem in this talk. Speculation.

Earth

- Below 100 km, ions and electrons collide with neutrals, not frozen to fieldlines
- Between 100 km and 160 km, electrons, but not ions, are frozen to fieldlines
 - Dynamo region, horizontal currents flow, but transport of plasma has no effect on electron densities
- Above 160 km, ions and electrons are frozen to fieldlines
 - When transport dominates photochemistry (>200 km?), ions and electrons move together along fieldlines
 - Ambipolar diffusion, no currents

Earth

- Two different theories for charged particle motion
 - Currents allowed, but gravity and pressure gradients neglected. Plasma transport does not change electron densities. Dynamo theory, sometimes with empirical electric field model
 - No currents allowed, plasma transport does affect electron densities. Ambipolar diffusion along fieldlines
- Works due to separation of regions
- Will fail on Mars

Conclusions

- Mars atmosphere is an integrated system. Lower and upper regions are not isolated.
- Chemistry (O / CO₂) affects energetics (CO₂ cooling), dynamics (winds) affect ionosphere (ion-drag), etc.
- Comparisons to Venus and Earth are helpful, but not always right.
- Effects of magnetic fields on ion/electron motion are an interesting problem.